

**PDF version of talk presented at the
DPS Annual Meeting, Oct 2012.**

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Cometary Delivery of Lunar Water: Transient Atmosphere Dynamics and Deposition Patterns

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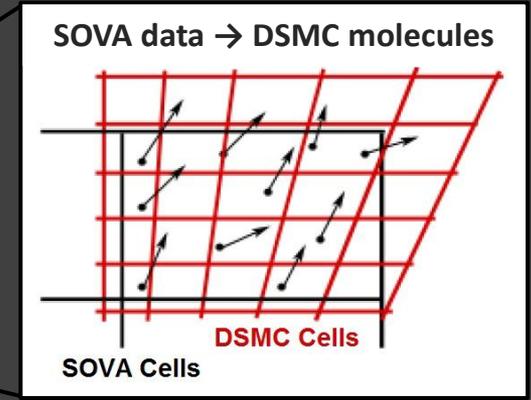
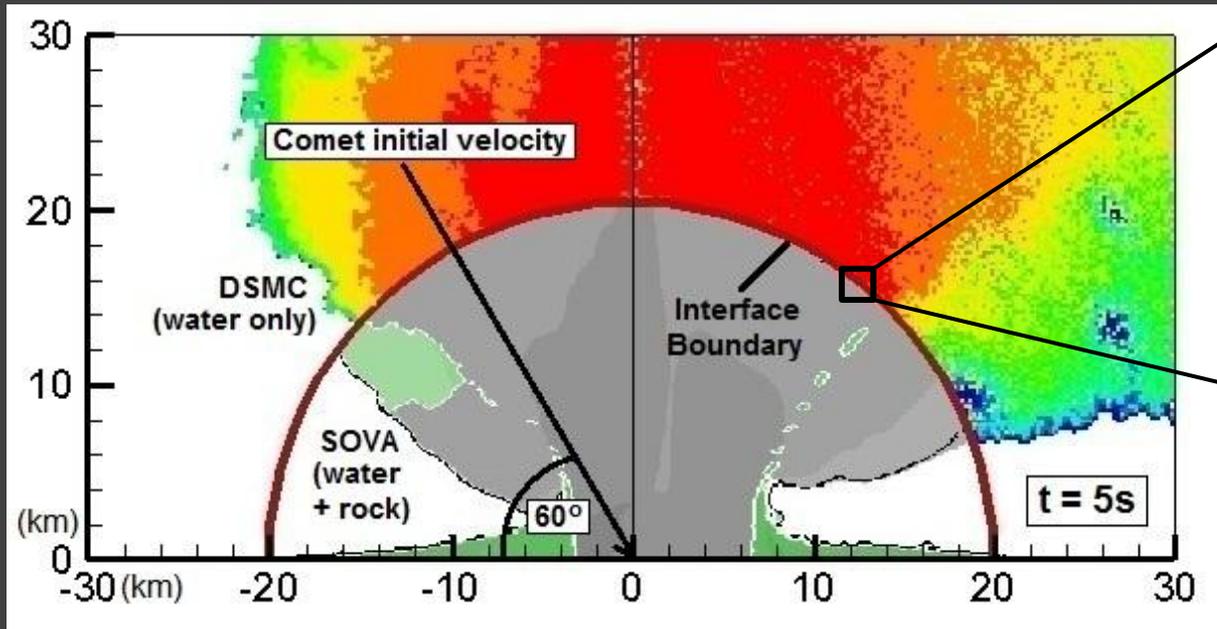
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Computations performed at the Texas Advanced Computing Center.*

- Observations by several missions- from Clementine (1994) and Lunar Prospector (1998) to LCROSS (2010) and others- suggest presence of lunar water, **particularly in craters with regions of permanent shadow (cold traps) at the poles.**
- Origins of lunar water?
 - Primordial water in the lunar interior.
 - Interaction of surface minerals with solar wind protons.
 - **Volatile-rich comet/meteorite impacts.**
- Questions:
 - How much water could comets have contributed to the lunar volatile inventory and deposited at cold traps?
 - How do impact parameters (angle, speed, location etc.) influence the quantity of cometary water retained?
 - Given a site indicative of a comet impact, what can we say about the associated volatile fallout?

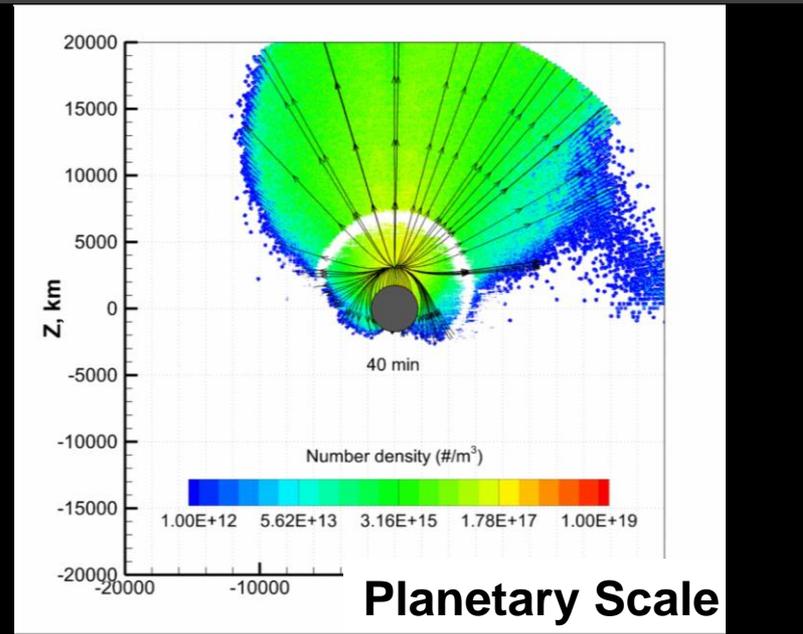
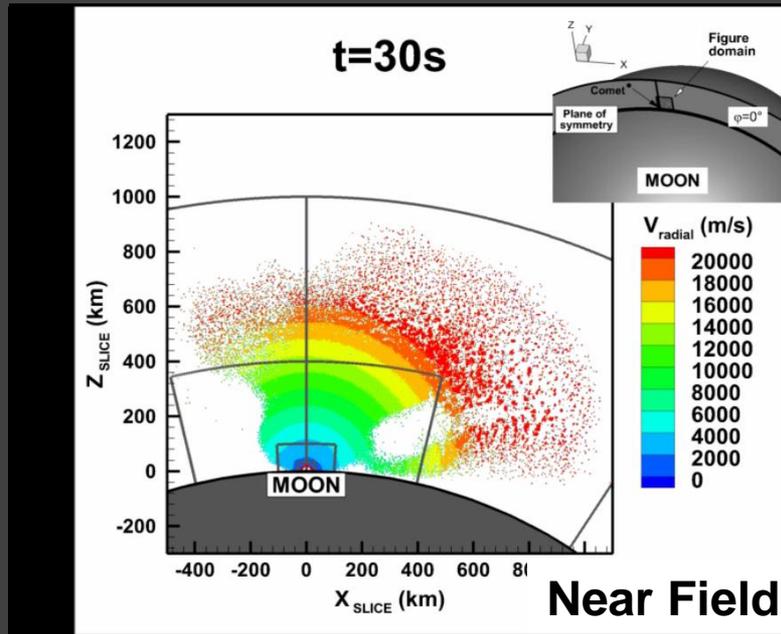
- Comet: pure H₂O ice, 1 km radius; target: dunitite.
- Varying impact angle (45°, 60°), speed (30 km/s, 20 km/s).
- Hybrid SOVA-DSMC method. DSMC code includes:
 - Surface temperature, dependent on distance from sub-solar point¹ and Moon's rotation.
 - 7 cold traps²: 1 NP (1257 km²) + 6 SP (4575 km²).
 - Temperature-dependent residence times³ for H₂O on H₂O ice matrix⁴.
 - Variable gravity.
 - Photodestruction probability⁵.

¹Butler, 1997; ²Elphic *et al.*, 2007 and Noda *et al.*, 2008; ³Langmuir (1916) and Frenkel (1924); ⁴Sandford and Allamandola, 1993; ⁵Huebner, 1992



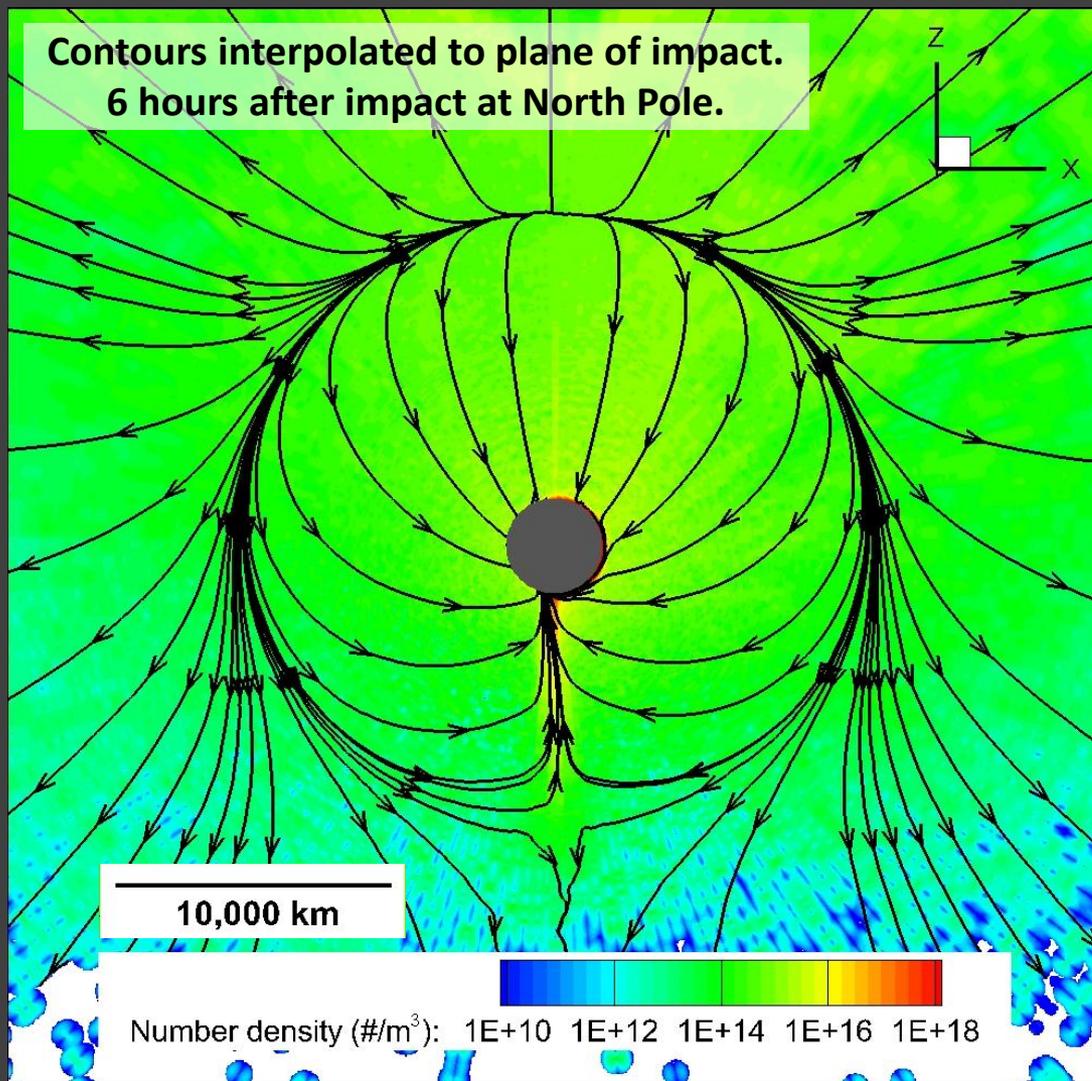
Supersonic speeds \therefore only SOVA \rightarrow DSMC coupling, none vice versa.

- **SOVA**: Simulates impact and hydrodynamic flow of relatively dense vaporized/molten comet and target material.
- **DSMC** (Direct Simulation Monte Carlo):
 - Simulates water vapor only. Particle based method- create, move, index, collide and sample 'molecules' (Bird, 1994).
 - Transition to rarefied but collisional expansion into vacuum.
 - DSMC is highly parallelizable.

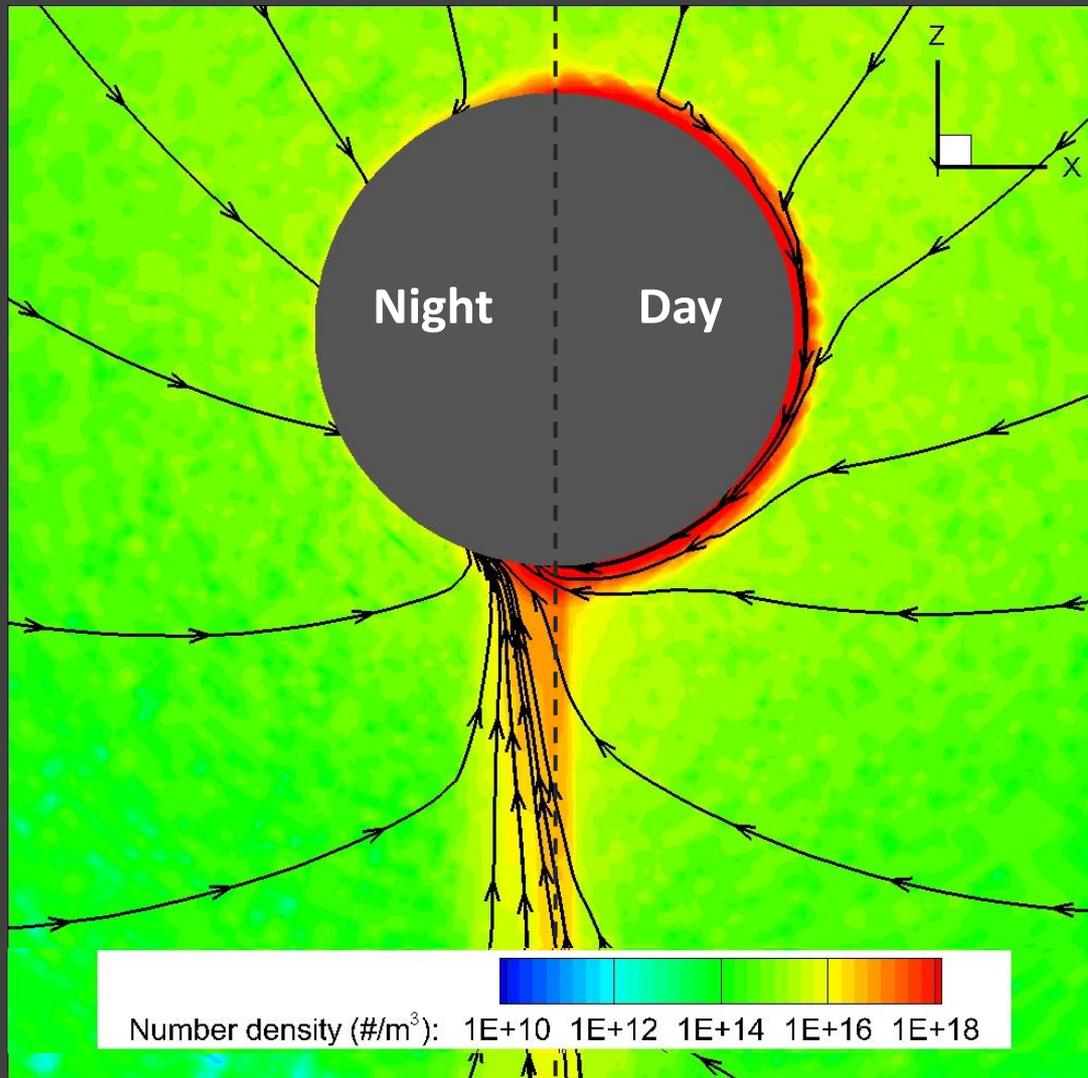


60°, 30 km/s impact. Contours interpolated to plane of impact.

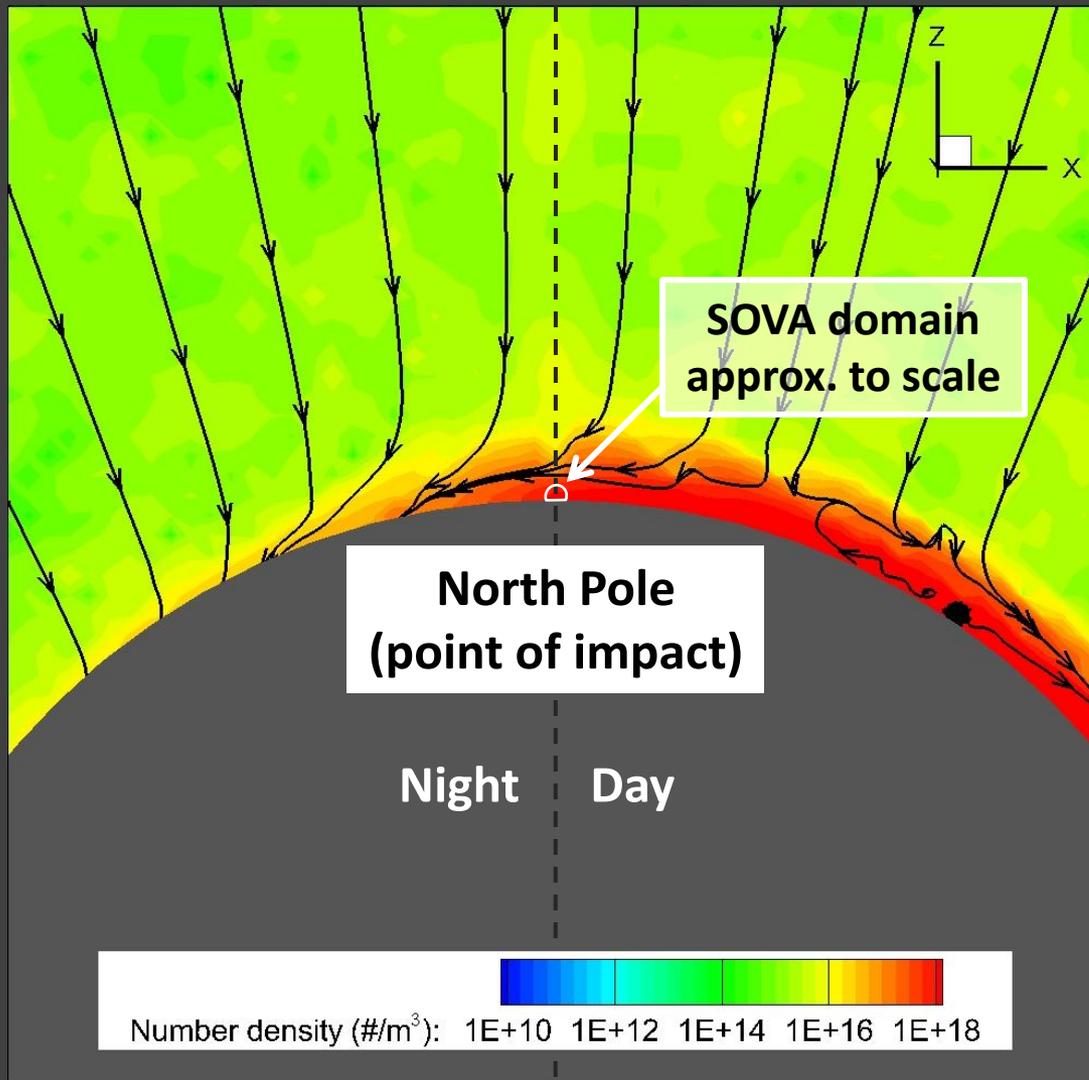
- Initial rapid radial expansion at velocities in excess of lunar escape velocity (2380 m/s).
- At later times, material enclosed by the white line travels below escape velocity, begins to fall back to surface.
- Later simulations neglect escaping vapor.



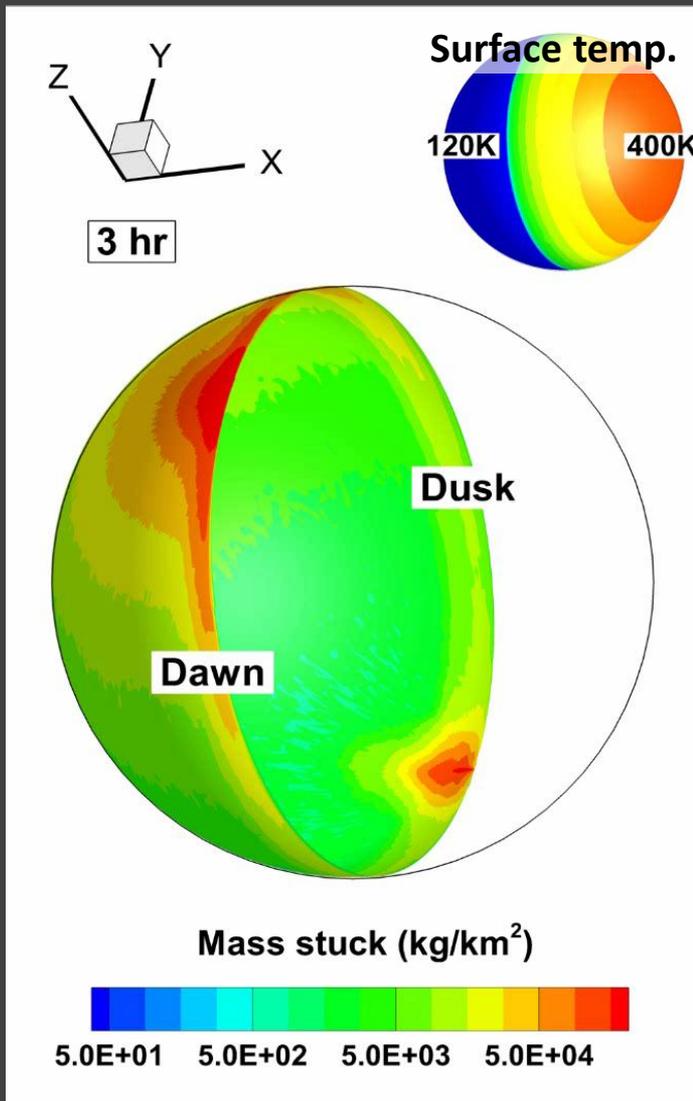
- Antipodal deposition. (Also seen in simulations of basin-forming impacts- Hood & Artemieva, 2008.)
- Ground-hugging day-side wind, driven by North-South pressure gradient. Reaches SP within 1-3 hours.



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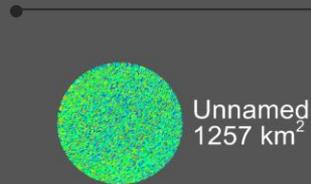
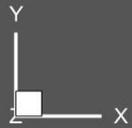


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- Ground-hugging day-side wind. driven by North-South pressure gradient. Reaches SP within 1-3 hours.
- Day-night pressure gradient \rightarrow region of reversed flow near point of impact (NP).



- Instantaneous snapshots of deposition patterns at 3 hour intervals. Cold traps not shown.
- Concentration of mass at poles due to day-side wind and antipodal convergence. $0(10^5)$ kg/km² \equiv ~ 0.1 mm ice.
- Concentration of mass at terminators due to day-night pressure gradient.
- Difference between dusk/dawn deposits due to fallback on to regolith with/without ice cover.

North Pole

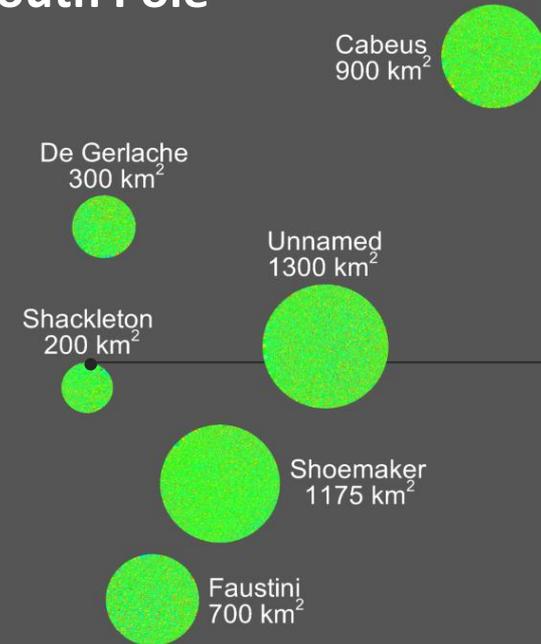
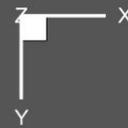


t=12 hours

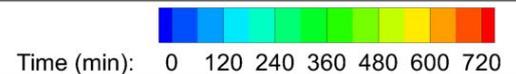


100 km

South Pole

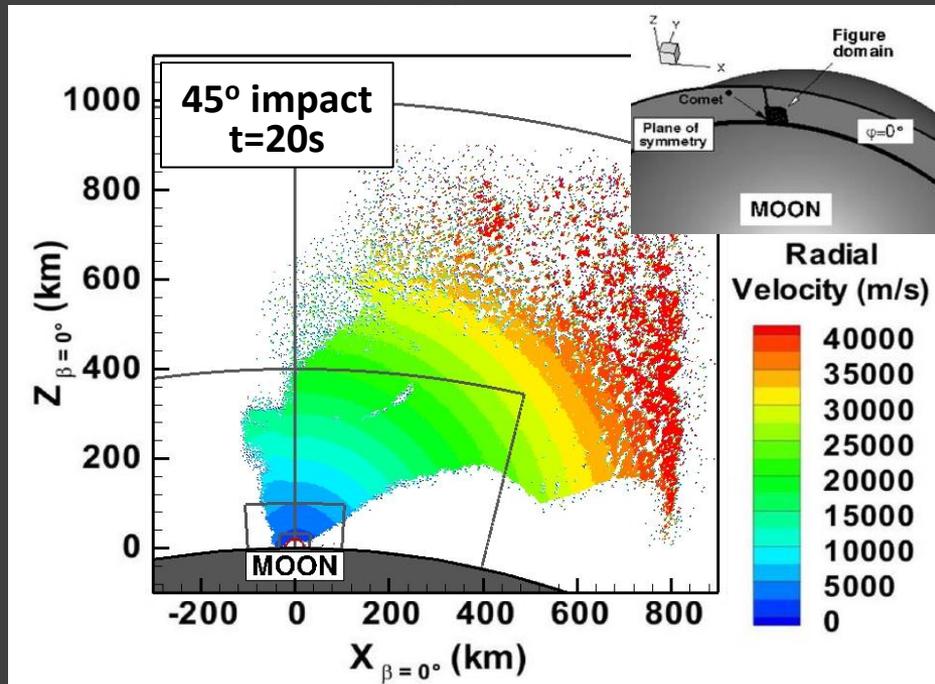
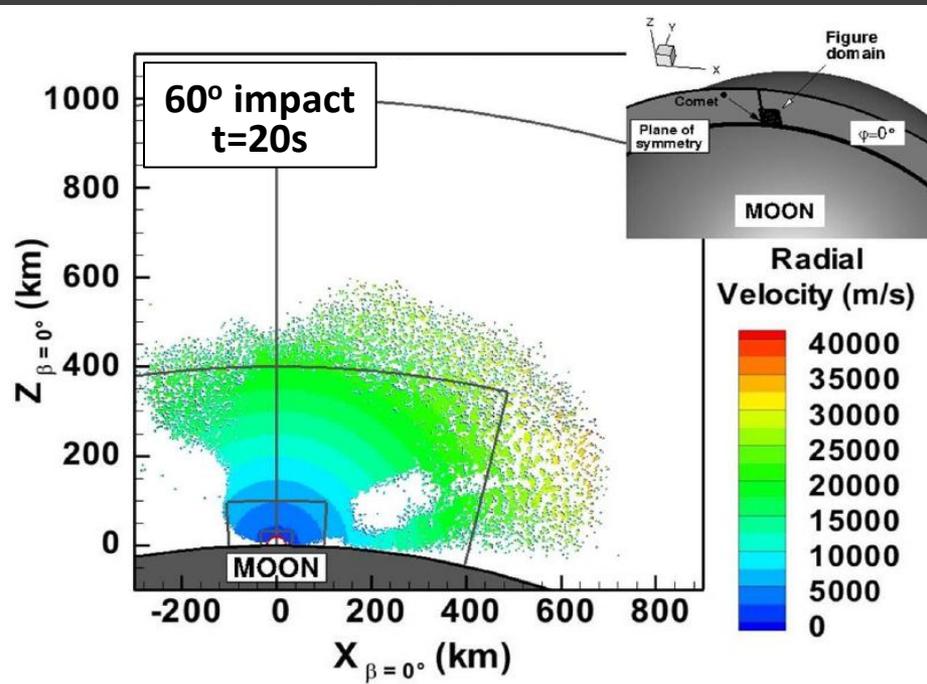


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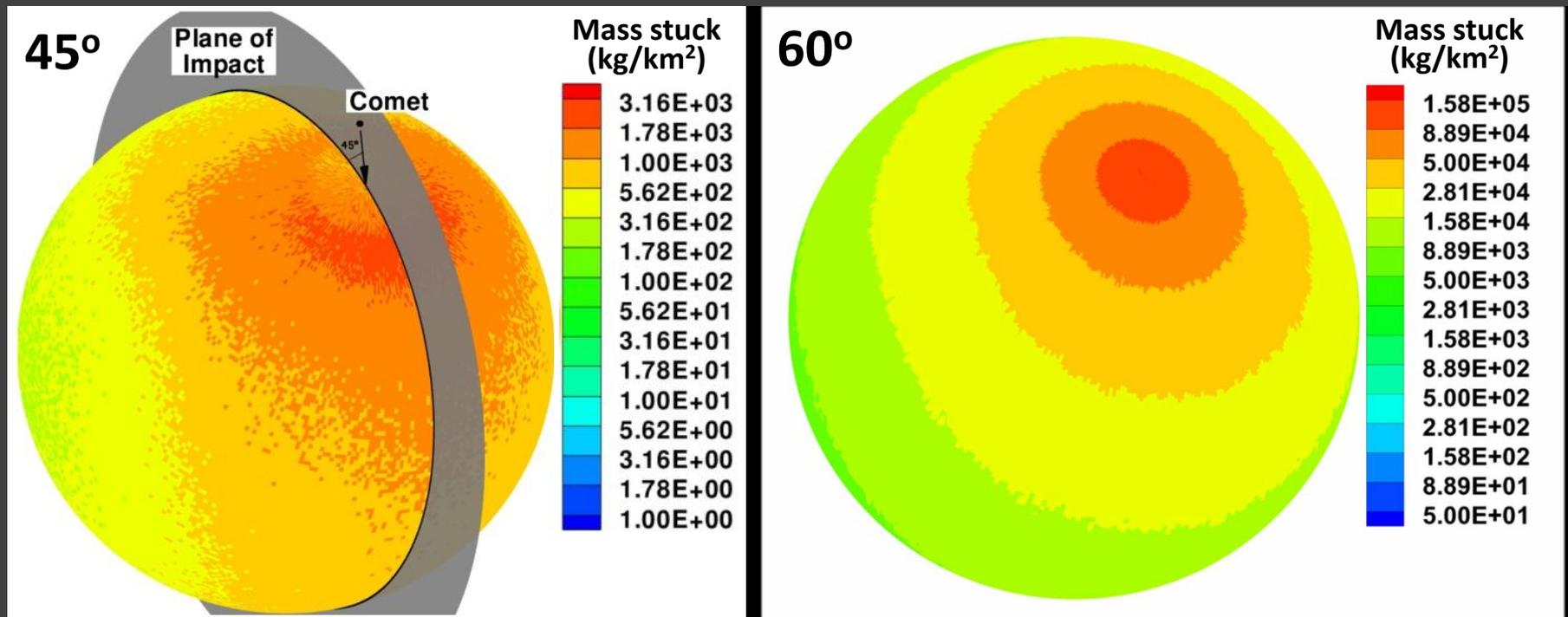


100 km

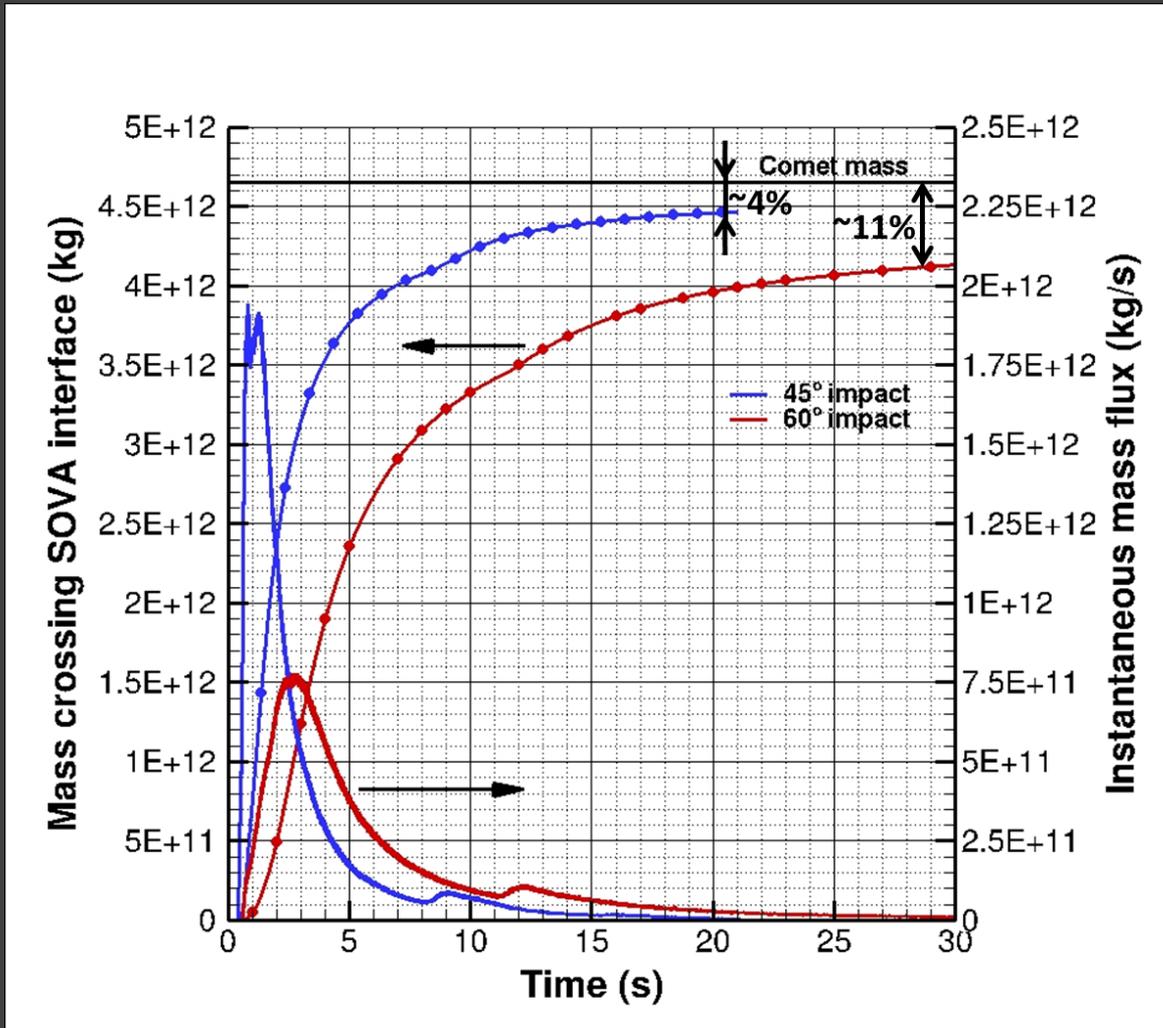
- Impact at North Pole → earlier deposits at North Pole cold trap.
- Antipodal convergence + day-side wind → later deposits at South Pole cold traps.
- Deposits colored by time of deposition.



- Less downrange focusing in the more vertical (60°) case → more symmetric (about point of impact) fallback.
- Lower plume kinetic energy in the more vertical case → less water escapes. (Consistent with findings of Pierazzo & Melosh, 2000 and Gisler *et al.*, 2006)



- Surface temperature fixed at 120K ∴ all water falling back sticks, no migration due to temperature variations. After 7 (Earth) days, most water has escaped or fallen back.
- Thicker (~50x) and more symmetric deposits in 60° case.



- SOVA simulations stopped when:
 - Outflow → subsonic.
 - Flux asymptotes → 0.
- Remaining water independently sublimated from a 'crater' centered at point of impact.
- ~10% of comet mass remains at the end of 60° SOVA run. ~1% for 45°.

Small amounts of water mixed with rock are neglected:
~1% comet mass for 60° run; ~3% for 45° run.

- **Stewart *et al.*, 2011 (Icarus 215, 1-16)** – long term (~6m) evolution and final results for a 45°, 30 km/s impact:
 - Similar retention rates for different impact locations - 0.14% of the comet mass is trapped $\equiv 1.12 \text{ kg/m}^2 \equiv \sim 1 \text{ mm ice}$ over all cold traps.
- 60° @ 30 km/s & 45° @ 20 km/s runs in progress.
 - Preliminary results suggest greater ultimate retention of water.
 - Dynamic transient atmosphere: day-side wind, antipodal deposition.
- Further considerations (not currently incorporated):
 - Radiative cooling of vapor.
 - Condensation (nucleation by dust/rock fragments?).
 - Optical shielding from photodestruction.
 - More complex treatment of “remaining water”.
 - Detailed topography/temperature maps.
 - Mechanisms acting after deposition in cold traps, for instance- space weathering (Crider & Vondrak, 2003), sub-surface migration of water molecules (Schorghofer & Taylor, 2007).